

scree. The corries in the granite of the Grampians (Fig. 37, see also Fig. 40), in the red sandstone mountains of the north-west, in the soft horizontal yellow sandstone of Orkney, in the gabbro of Skye, in the Silurian grits and shales of the southern counties, may be cited as examples of this diversity of type. All the Scottish corries have been occupied by glaciers: hence their bottoms are generally well ice-worn or strewn over with moraine stuff. Not unfrequently also a small tarn fills up the bottom, ponded back by a moraine (Figs. 37, 71). It is in these localities that we can best observe the last relics left by the glaciers that once overspread the country.

The proofs of enormous and comparatively rapid disintegration, which the loftier mountain solitudes present to us, appeal powerfully to the imagination. They compel the admission that, during a period of sufficient vastness, the sub-aërial agents of waste could carve out glens without the aid of subterranean convulsions. In the scenery around Ben Nevis, there is no evidence of any open fissures or dislocations to which the glens could have owed their origin. On the contrary, these deep, narrow valleys are crossed at their upper ends by the precipitous junction of their sides, and the connecting wall of granite, though seamed and cracked by the weather, shows no trace of subterranean fracture. Natural, therefore, as the impulse undoubtedly is, in presence of such impressive scenes, to refer these profound concavities to the sundering of the sides of fissures, a little examination of the ground and reflection on the conditions of the problem will assure us that no subterranean dislocation will explain what we see. The two deep glens to the east of Ben Nevis are separated by the narrow granite ridge of Carn Mor Dearg already described; if one of them had been opened by a rending