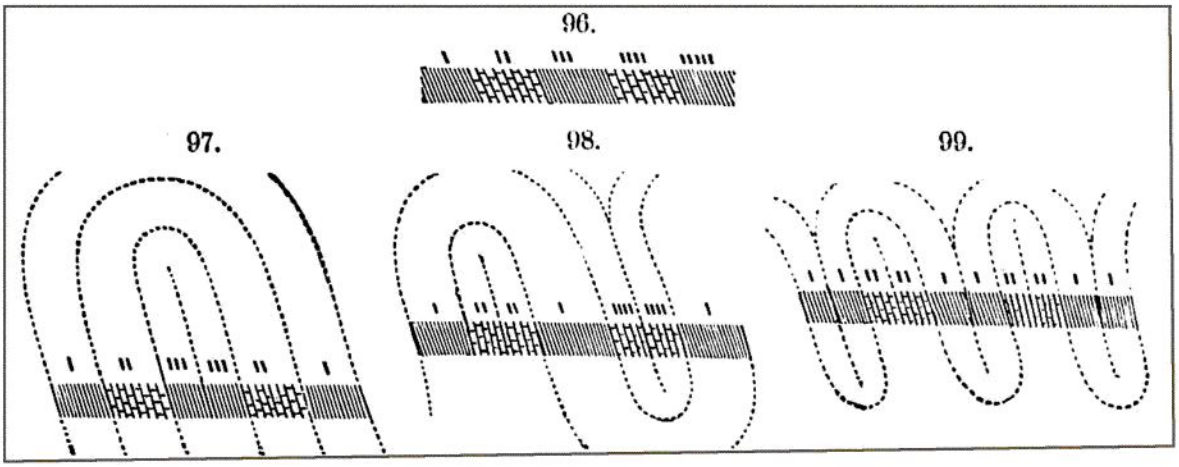
Fig. 96 below represents a section of alternating belts of limestone and schist, numbered I to IIIII, to be interpreted.

It may be that each belt, I, II, III, IIII, IIIII, is an independent stratum, alike in dip, with IIIII the highest in the series. This is the simplest explanation. But there may be flexures, and Figs. 97, 98, 99 represent some of the possible methods of interpretation. By comparing each with Fig. 96, the relations may be studied out.



In a region of flexed rocks the same bed, as the illustrations show, may come many times to the surface; and it is therefore easy for the observer to be deceived in such regions as to the number of independent beds. The covering of soil adds greatly to the difficulty, as the following figures illustrate. When the rock in a region of high dips is

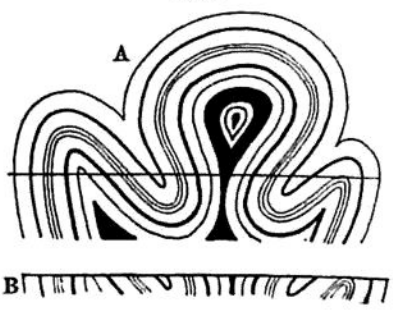
100.

101.

simply a slate or shale, with no associated stratum of permanent horizon, it is almost impossible to decide as to flexures. Such beds bend easily, and may be full of flexures, and yet none may be apparent.

Sometimes an anticlinal flexure has the dips of a synclinal, as in the central part of Fig. 102 A. If worn down to a plane (Fig. 102 B), the dips along the center would seem to be good evidence of a syncline. Such *fan-shaped* folds are common on a small scale in schists, and occasionally they may occur on a scale of mountain magnitude. The facts at Mont Blanc in the Alps are explained on the idea of such a fold.

102.



To reach positive conclusions among the possible explanations, the beds or strata must be carefully compared, and also the sides and middle of the several strata, as to texture and all other differences. Besides, search should be made for outcrops that exhibit the limestone and schist in *broad* anticlinals or synclinals, as in the following cases. In Mount Washington and Greylock of the Taconic range on the boundary of western New England, the beds dip at the north end of the mountain mass, nearly