stones, however, contain argillaceous, calcareous, or ferruginous concretions which weather more rapidly than the surrounding rock, and cause it to assume a honeycombed surface; others are full of a diffused cement (clay, lime, iron) the decay of which makes the rock crumble down into sand. In sandstones, as indeed in most stratified rocks, there is a tendency toward more rapid weathering along the planes of stratification, so that the stratified structure is brought out very clearly on natural cliffs (Fig. 92). In many ferruginous sandstones and clay ironstones, successive yellow or brown zones or shells may be traced inward from the surface, frequently due to changes of the ferrous carbonate into limonite, the interior remaining still fresh. In many prismatic massive rocks (basalt, diorite, etc.) segments of the prisms weather into spheroids, in which successive weathered rings form crusts like the concentric coats of an onion (Figs. 93, 94). Where one of these rocks has been intruded as a dike, it sometimes decomposes to a considerable depth into a mass of brown ferruginous balls in a surrounding sandy matrix—the whole having at first a resemblance to a conglomerate made of rolled and transported fragments (Fig. 95).

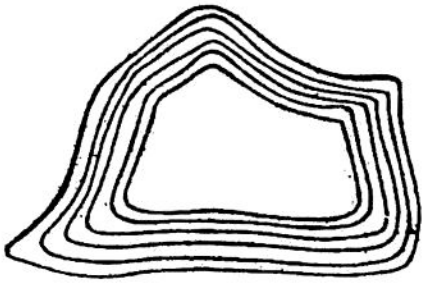

Fig. 93.—Rings of Weathering.

No rock presents greater variety of weathering than granite. Some remarkably durable kinds only yield slowly at the edges of the joints, the separated masses gradually assuming the form of rounded blocks like water-worn bowlders. Other kinds decompose to a depth of 50 feet or more, and can be dug out with a spade. In Cornwall and Devon, the kaolin from the rotted granite, largely extracted for pot-