

and condensation may be thus considered as mutually dependent; so that one process cannot take place without the other. For this reason, in the great expanse of nature, these two processes oscillate or fluctuate about the point of equilibrium, within certain limits which are never passed; and which limits, though subject to countless anomalies, in general, decrease from the Equator toward the Poles.

With respect to the temperature which constitutes the point of equilibrium; in an atmosphere of vapour, that point would, of course, be the maximum point of saturation. But in a mixed atmosphere of vapour and air, like that of our globe, the point of equilibrium cannot be the point of utmost saturation, but must be that *inferior point of saturation* formerly described, as being determined by the temperature of the predominant air. Thus at the Equator, where the mean temperature at the level of the sea is about 80° , the mean point of saturation will, according to our former estimate, be 61° ; while in London, where the mean annual temperature is about $49\frac{1}{2}^{\circ}$, the mean point of saturation, (or the *dew point*, as it is termed,) has been fixed by Mr. Daniell at $44\frac{1}{2}^{\circ}$. In temperate climates, the mean point of saturation, at any particular place, varies with the seasons from day to day, being higher in summer than in winter. During