frequent occurrence of extensive beds of calcareous conglomerate connected with the newer magnesian limestone. (C.)*

(*a*) *Chemical and external Characters.* This limestone contains about 20 per cent of magnesia, and prevails considerably in England.

Analysis by Smithson Tennant, Esq. of the stone of York Minster.

Carbonic acid	47.00
Lime	33.24
Magnesia	19.36
Iron and clay	0.40
	100.

That of Westminster hall contains about 2 per cent. less of magnesia. That of Denton near the Tees, consists, according to the analysis of the Rev. J. Holme, about the same proportion of the latter, with a very small proportion of bitumen, and less than 1 per cent. of water. (G. Notes.)

A magnesian lime from Eldon, analysed by Sir H. Davy, yielded Carbonate of lime 52, Carbonate of magnesia 45.2, Iron 1.1, Residuum 1.7. Another from Denton, not far from the Tees in Durham, analysed by the Rev. J. Holme, yielded 11 per cent. more carbonate of lime, and 11 per cent. less of carbonate of magnesia. (G. T. vol. 4. p. 7.)

It differs from common limestone in external character, in having generally a granular sandy structure, a glimmering lustre, and a yellow colour; and in the course of the range from Nottingham northwards, its surface in many places is covered by a poor herbage, uncommon to limestone, and attributable to the magnesia it contains, which is known to be unfavourable to vegetation. It is associated with a conglomerate limestone. This conglomerate often exhibits very distinct fragments of the oldest mountain limestone, passing gradually from an aggregate compound of very large pebbles of this rock, to one in which the grains are so small as scarcely to be distinguishable.

In a quarry at Hartlepool on the coast of Northumberland, is a stratum of hard white oolite, the grains composing it being about the size of a mustard seed.

* Though the quantity of magnesia contained has been considered as in some degree characterising this formation, yet it is a character which, taken singly, cannot be relied upon, since magnesia in considerable quantity has likewise been found in some of the chalk of France, in some of our own oolites, in the mountain or carboniferous limestone, in that associated with transition rocks, and, as need not be stated, in primitive dolomite: we have not however ventured to change a name generally received. (C.)