

distance from the snow-field to the foot of the glacier looks not much more than one English mile—at least it is but short compared with the rapidity of descent. Hence the glacier is steep, and in some places much crevassed. Issuing from the upper snow in a steep, broken, and jagged slope of blue ice, it descends by a series of steps, till, getting compacted again in the valley below, it passes into a solid, firm glacier, with a tolerably smooth surface, forming a declivity of 12° or 15° (Fig. 10). At a point about half a mile or less from the foot of the glacier the valley suddenly contracts, and the glacier, much narrowed and compressed, tumbles over a second steep declivity in a mass of broken ice. The crevasses speedily unite, and after another descent of 300 or 400 yards at an angle of 25° , the glacier comes to an end. At the point where the strangulation takes place the glacier lies in a kind of basin, of which the lower lip presents proofs of the most intense erosion. On the western bank, in particular, a mass of the mountain side which projects into the ice has been ground away, and shows plainly enough, by its form and striæ, that the glacier, ascending from the basin, has climbed up and over this barrier, so as to tumble down its northern or seaward side.

The course of this little glacier is now too short to admit of the formation of moraines. Yet there are large heaps of rubbish and enormous masses of rock scattered over the valley below; and the moraine at Fondalen is a further proof that, when the ice formerly filled the valley, its surface received abundant detritus from the mountain slopes on either side.

Opposite Fondalen, the Holands Fjord, passing through a deep and narrow channel on its northern bank, trends in an east-north-easterly direction; but just before taking this course it sends eastward a bay which terminates at the