rides the moraine heaps, ploughing them and pushing them over (Fig. 12). On the west side of this prolongation of the glacier the ice is separated from the moraine mound by a small lake, of which the surplus waters find their way seaward by cutting through the moraine. Like many lakes still existing in Britain, this sheet of water is formed by the dam of rubbish thrown down by the glacier across the valley. It is full of fragments of ice, which break off from the parent mass, and float across to the north or lower side, where they strand on the moraine heaps, and gradually melt away. The smaller pieces, however, often find their way into the stream by which the lake discharges itself, and are then carried down into the fjord. From the mean of several observations taken with the aneroid, I estimated the surface of this lake to be about 25 feet above the level of high water in the 1jord. We had no means of measuring its depth, yet, from the slope of the glacier, it may be inferred that the bottom of the ice is probably lower than the level of the sea.

Proofs that the glacier was once much larger than it is now may be well seen on the west side of the valley, a little above the lake. The shelving slopes of the mountain for several hundred feet upward have been shorn smooth, grooved, and striated, and every polished hummock of rock is loaded with huge fragments of stone and heaps of earth and angular rubbish. Here, as at every glacier we visited, the glaciation of the rocks, down to the minutest detail, was exactly similar to that of the coast and outer islets, as well as to that of the Scottish glens and sea-lochs.

But the feature which most interested us was the relation of this large glacier of Fondalen to the marine deposits of the locality. The foregoing sketch-map (Fig. 11) shows that the high terrace so marked along the sides of the