

are, as a rule, deep, sometimes much deeper than the sea outside. But the sea, as we have seen, cannot scoop out deep hollows; its erosive power is confined to its upper part, and its tendency is to plane down its shores. Instead of excavating a fjord, it can only level the rocks that are lashed by its breakers. If the land-valley has been hollowed out by sub-aërial erosion, a like origin must be assigned to the continuation of the same valley under the sea-level. The sea-lochs of the west coast are assuredly not inlets cut out by the waves. They are old glens that have been submerged beneath the sea. They prove that the west side of the island has subsided in a comparatively recent geological period, and that the tides now ebb and flow where of old there was the murmur of brooks and waterfalls.

If this view of the origin of the western sea-lochs be correct, it is natural to expect that traces of different stages of the submergence should be found: that, as the downward movement of the land went on, some lake-basins in the valleys should have been carried far down beneath the surface of the sea; that others at a higher level should have sunk a shorter distance, and that others should have barely escaped when they had approached within a few yards of the sea-level. Examples of these various steps in the process will occur to those who are familiar with the western coasts. Of the first, Loch Fyne is a notable illustration, as it deepens a little south of Tarbert into a basin 624 feet below the surface of the loch, and shallows northward and southward. If this great fjord were now a land-valley, that basin would be filled with a lake. Of the second stage, Loch Etive forms a good example. This inlet narrows at Connal Ferry, and across the straitened part runs a reef of rocks covered at high water, but